Identifying the Earliest Neolithic Settlements in the Southeastern Balkans
Laurent Lespez, Zoï Tsirtsoni, Pascal Darcque, Dimitra Malamidou, Koukouli-Chrysanthaki Chaido, Arthur Glais

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Going West?
The Dissemination of Neolithic Innovations between the Bosphorus and the Carpathians

Proceedings of the EAA Conference, Istanbul, 11 September 2014

Edited by Agathe Reingruber, Zoï Tsirtsoni, Petranka Nedelcheva

Going West? uses the latest data to question how the Neolithic way of life was diffused from the Near East to Europe via Anatolia. The transformations of the 7th millennium BCE in western Anatolia undoubtedly had a significant impact on the neighbouring regions of southeast Europe. Yet the nature, pace and trajectory of this impact needs still to be clarified. Archaeologists previously searched for similarities in prehistoric, especially Early Neolithic material cultures on both sides of the Sea of Marmara. Recent research shows that although the isthmus of the Dardanelles and the Bosphorus connect Asia Minor and the eastern Balkans, they apparently did not serve as passageways for the dissemination of Neolithic innovations. Instead, the first permanent settlements are situated near the Aegean coast of Thrace and Macedonia, often occurring close to the mouths of big rivers in secluded bays. The courses and the valleys of rivers such as the Maritsa, Strymon and Axios were perfect corridors for contact and exchange not only in a south-north direction but also the other way round. Using previous studies as a basis for fresh research, this volume presents exciting new viewpoints by analyzing recently discovered materials and by applying modern research methods of interdisciplinary investigations.

The seventeen authors of this book have dedicated their research to a renewed evaluation of an old problem: namely, the question of how the complex transformations at the transition from the Mesolithic to the Neolithic can be explained. They have focused their studies on the vast area of the eastern Balkans and the Pontic region between the Bosphorus and the rivers Strymon, Danube and Dniestr. Going West? thus offers an overview of the current state of research concerning the Neolithisation of these areas, considering varied viewpoints and also providing useful starting points for future investigations.

Agathe Reingruber, researcher at the Freie Universität Berlin, is specialized on topics related to the Neolithic and Chalcolithic of southeast Europe (Greece, Turkey, Romania). She is currently running a project in northeastern Thessaly focusing on population dynamics.

Zoï Tsirtsoni, researcher at the French National Centre of Scientific Research (CNRS, Laboratory Archéologies et Sciences de l’Antiquité, Nanterre), is specialized on the Neolithic, Chalcolithic and Bronze Age periods in the Aegean and southern Balkans. She is co-director of the Greek-French research project at the multilayer (tell) settlement of Dikili Tash in northern Greece.

Petranka Nedelcheva, Assistant Professor at the New Bulgarian University (Sofia), is a lithics specialist for the Neolithic, Chalcolithic and Bronze Age periods in southeast Europe, western Anatolia and the Caucasus. She participates in several projects in Greece, Turkey, Romania and Georgia.
Identifying the Earliest Neolithic Settlements in the Southeastern Balkans
Methodological Considerations Based on the Recent Geoarchaeological Investigations at Dikili Tash (Greek Eastern Macedonia)

Laurent Lespez, Zoi Tsirtsoni, Pascal Darcque, Dimitra Malamidou, Haïdo Koukouli-Chryssanthaki and Arthur Glais

Introduction
For many decades the difficulty to locate sites with permanent installations datable to the 7th millennium bc in the area immediately west of the Bosporus (Thrace and eastern Macedonia: Figure 4.1) has nourished the idea that this part of Europe entered the sedentary way of life later than its neighbours, namely later than the rest of Greece (Demoule & Perlès, 1993: 365, 388; Andreou et al., 1996: 586; Perlès, 2001: 59–60). This premise, which has been maintained even after the discovery of such early layers at Hoca Çeşme, at the mouth of the Maritsa/Meriç/Evros River (Özdoğan, 1993, 1997; see also Karul, this volume), has been further used in several models about the Neolithic expansion throughout Europe. The latter put forward the existence of presumed stopping places that coincided with the topographical, ecological or agro-ecological barriers met by the first settlers (Van Andel & Runnels, 1995: 494–8; Guilaine, 2001; Rasse, 2008; Özdoğan, 2011a, 2011b; Guilaine, 2012: 10), and acted as ‘centres of renewed expansion’ (Bocquet-Appel et al., 2009; see also syntheses by Reingruber, 2011; Vander Linden, 2011). In this chapter we present evidence that contradicts this claim, while at the same time raising questions about the limits to the conclusions based on standard archaeological investigations (survey and digging) and suggesting supplementary tools for overcoming the biases.

State of the Questions
In the specific area of the Greek eastern Macedonia, the delay of the start of the Neolithic seemed even longer than in Thrace, as no site was known, not only for the late 7th millennium bc, but for the early 6th millennium bc as well. Indeed, with the exception of a few dubious white-on-red painted sherds from Toumba Serron (Grammenos & Fotiadis, 1980: 17, 20–3), no site had yielded surface finds that could be attributed to the Early Neolithic period as defined for the Aegean (Alram-Stern, 1996: 88–9; Andreou et al., 1996: 538; Treuil et al., 2008: 58–9). All recognisable artefacts from surface collections were assigned to phases contemporaneous with the Thessalian ‘pre-Dimini’ stages (Tsangli, Arapi), or later (Grammenos & Fotiadis, 1980; Grammenos, 1991: 99–105, 120–6; Koukouli et al., 2008: 397–403). Furthermore, the unique major site where excavation reached the virgin soil, the tell of Sitagroi, seemed to start towards 5500 cal bc (Renfrew et al., 1986: 175–81, 27: table 2.1), that is at a stage parallel to the very end of the Thessalian Middle Neolithic Sesklo culture and the transition to the Late Neolithic. This date was sustained also by later excavations at the neighbouring tell of Dimitra (Grammenos, 1991: 45–6) and at the insular site of Limenaria on Thasos (Papadopoulos & Malamidou, 2012: 38–9, 282), which also reached the natural soil (see also Papadopoulos, 2009: 51–2).

Both of these arguments – the argument of surface finds and that of the lowest excavated levels – should be considered, however, with great caution. The idea that the artefacts scattered on the ground at the top of an ancient site would represent the entirety of its occupation sequence is indeed highly contested (Bintliff et al., 1999), despite some objections about the general character of the phenomenon (Mee & Cavanagh, 2000; Davis, 2004). The representation, in particular, of the earliest occupation levels in the surface material seems even more problematic at sites with a strong sedimentation, whether the latter results from human activities (multi-layered settlements) or from natural factors (alluviation).1 In Greek eastern Macedonia, like in the

1A good example has been recently presented by Dr. St. Chryssoulaki at the International Conference about Athens and Attica in Prehistory (Athens, ASCA, 27–31 May 2015): a well-preserved Early Bronze Age settlement discovered near the riverbed of Kifissos, beneath 7 m of alluvial deposits. The excavation did not manage to continue below this level.
rest of southeast Europe, most Neolithic settlements identified so far are tell sites – a situation that reflects prehistoric settlement pattern as well as the orientation of archaeological research (Gaul, 1948; Grammenos, 1996; Bailey, 2000). On the other hand, the large plains of eastern Macedonia have undergone important alluvial, lake and marshy sedimentation during the Holocene. As demonstrated, in the lower Strymon valley, Late Neolithic sites are covered by more than 7 m of sedimentation (Lespez et al., 2014). We can therefore assume that many flat or low-elevated settlements have been neglected in non-systematic surface surveys, or have been buried on the edge of the alluvial plains between the Bosporus and western Macedonia.

In the case of tell sites, we are also facing the problem of the extent of the lowest excavated levels. In most cases indeed, these are only explored in just one or two very small trenches – again because of the great thickness of the overlying deposits, which does not allow extensive exposure, unless by accident (public works, looting, etc.). Thus, our knowledge about the presumed start of occupation at Sitagroi towards 5500 cal bc relies on the results of a unique 9-square-metre trench opened approximately in the centre of the tell (deep sounding ZA: Renfrew et al., 1986: 17–8 and Fig. 2.2); this applies much the same for Dimitra and Limenaria (supra). Considering that the first installations were probably smaller than the later Neolithic and Bronze Age settlements (cf. for instance Knossos: Evans, 1971: pl. VI), one might presume that they could also well be located at some other part of the perimeter of the future tell, or even beyond it.

This claim was supported by results from recent investigations at the mound of Krový in the neighbouring area of Aegae Thrace, using core-drills into the natural soil (Ammerman et al., 2008). The earliest radiocarbon-dated level here (first half of 6th millennium bc at about 40 cm above the natural palaeosol) was recorded not in the centre, but near the eastern edge of the mound (core 2: see Ammerman et al., 2008: 143–4, Fig. 2), although it is not entirely clear whether this was due to a lack of adequate organic samples or a lack of deposits in the other cores.

Our own research in the years after 2010 on and around the tell settlement of Dikili Tash, in the southeastern part of the Drama-Philippi plain, goes further in this direction; it proves indeed that part of the current archaeological picture and the discourse that accompanies it are seriously biased by taphonomical problems. Using this experience as a starting point, we plead here for a closer collaboration between archaeologists and geomorphologists, and we propose an entire

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**Figure 4.1.** Map of the Southeastern Balkans With Neolithic Sites of the 7th and Early 6th Millennium bc. Dikili Tash Is Shown In Bold.
range of analytical procedures (in the field and in the laboratory), which could help remedy this problem. More than just filling the gap in the regional distribution maps, this knowledge should provide the necessary framework for any discussions about the conditions and the pace of the spread of the Neolithic way of life into the Balkans. Some of the results have been already presented elsewhere (Darcque, 2013: 69–72; Lespez et al., 2013; Darcque et al., 2014), but others are discussed here for the first time.

**THE FRAMEWORK OF RESEARCH AT DIKILI TASH**

The site of Dikili Tash is rather well known in the archaeological literature; therefore, it does not seem necessary to proceed here with a detailed presentation (a full, up-to-date bibliography is given at www.dikilitash.fr; see also references cited above and below). Three points should be recalled however, which are relevant to our topic.

The first point is the topographical situation of the site: at the foot of a low mountain range (Lekani), originally covered by forests, right next to an important freshwater spring and close to the lowest parts of an intra-mountainous floodplain covered by water (swamp) during the entire Holocene period (Treuil, 1992: 3–11; Lespez et al., 2000: 415–7; Lespez, 2008: 260–7). This privileged location had obviously favoured human settlement in all periods from the mid-6th millennium BC onwards, that is the periods attested by levels excavated archaeologically (Darcque & Tsirtsoni, 2010), and legitimated the assumption that the site might have attracted also earlier populations, including those first adopting a sedentary way of life.

The second point concerns the state of archaeological research prior to the investigations that we discuss here (i.e. prior to 2010). Excavations at the tell, whose top lies today at about 15 m from the surrounding surface, proceeded generally in a step-like manner, with lower trenches starting approximately at the point where those located higher – in terms of both hypsometer and relative chronology – stopped. Thus there existed no unique ‘deep’ stratigraphic trench like that in Sitagroi (supra), but several ‘medium-depth’ trenches starting from various altitudes and cutting through different parts of the tell’s deposits. The trench that seemed to get closer to the start of the settlement was sector 1, pursuing the older ‘AA28 trench’ on the southern slope of the tell (Treuil, 1992: 15, 20–1; Darcque & Tsirtsoni, 2010: 59). The excavation had stopped there at about 4 m above the estimated natural soil level (based on a series of core-drills made in and around the tell in 1993: Lespez et al., 2000: 417–21), inside a destruction level of the early LN I period (c.5400/5300 cal BC). Another trench made lower in the southern slope (sector 2) got even closer to the virgin soil in terms of hypsometer, but the excavated deposits were mostly composed of later features and colluvia, suggesting that here we were at the limits or even outside of the settlement properly speaking (Darcque et al., 1990: 877; Darcque et al., 1992: 715; Darcque et al., 2009: 532–33).

The third and last point deals with evidence, or rather non-evidence, from surface finds. Indeed, one century after the site’s first identification by surface collection (Welch, 1918–1919), followed by fifty years of agricultural exploitation and another fifty years of systematic archaeological investigation (Darcque, 2013; Treuil, 2014), with all the circulation, deposition and reworking of sediments that these activities involve, we knew no more than two (!) potentially Early to Middle Neolithic sherds from Dikili Tash – one with white-on-red painted decoration assigned to the horizon of Karanovo I-Kremikovci (Deshayes, 1970: 808 and Fig. 19), and one with channelled decoration assigned to Karanovo II (Deshayes, 1973: 472–3, Fig. 13). This absence of earlier material contrasted sharply with the overall estimated tell’s sequence as was established from the two intra-site drills of 1993 (supra): the latter not only showed, as we said, that there existed several meters of unexplored archaeological deposits at the base of the tell, but also provided a series of very promising radiocarbon dates in the early 6th millennium BC, and even one from the mid-7th millennium BC (mentioned by Lespez et al., 2013: 33; see also infra).

Therefore, the two main questions that needed to be answered concerned:

1. The date of the first settlement at Dikili Tash;
2. The precise location of the earliest levels; concretely, we wished to know whether we could reach them by continuing down the sector 1 (i.e. by excavating the 4 m left to the virgin soil, which would be the most economical solution in terms of both time and effort), or if we had to look elsewhere.

In order to answer to these questions, it was decided to make a new series of core-drillings, not only to the south but also at the other sides of the tell, east, northeast and northwest, the last two being until then ‘terrae incognitae’.

**THE 2010–2013 GEOARCHAEOLOGICAL RESEARCH: METHODOLOGY**

The four main drillings made in 2010 (C1 to C4, depth from 6 to 10 m) were set in the tell as shown in Figure 4.2. Two smaller drillings (depth 1 to 3 m)
Figure 4.2. Map of the Core Drillings and Interpretation of the Settlement Dynamics During the Early Stages of the Neolithic. 1. First phase of Occupation (c.6500–6200 BC) on top of the Holocene Palaeosol. 2. Possible Extension of the Initial Settlement. 3. Limit of the Marshy and Alluvial Deposits During the Cold and Wet Period Around 6100–6200 BC. 4. Possible Extension of the Settlement During the Wet Phase, Assessed Only in core C10.

were made inside the limits of sector 2 (supra), in order to clarify the depositional sequence there. They were completed in 2012 by four additional intra-site cores (C7 to C10), whose aim was to further refine the overall occupational sequence and suggest more firmly the place where we should dig in order to reach the earliest levels with the less possible effort. Finally, a series of six drillings (Dik1 to Dik11, depth from 3 to 5.5 m) was made in 2013 in the immediate surroundings of the tell, in order to investigate further the relationship between the settlement and the nearby wet zones, namely the extended Philippi marsh to the south and the small pond formed by the Dikili Tash water source to the north; further sampling in these off-site areas was made in 2014.

All drillings were realised with a hand-driven percussion device (Cobra TT) equipped with gouges of 60 mm diameter and 1 m long; work in 2012 and 2013 was facilitated by the use of a small hydraulic extractor (Figure 4.3). This kind of device seems better adapted for coring in anthropogenic layers than devices driven by heavy machinery, which tend to compact too much
the relatively loose sediments and to disturb more deeply the deposits.

Practically all intra-site cores and most of those made off-site (Dik6, Dik7, Dik10, Dik11) were taken with the open gouge method. In this method, the sediment is carefully extracted from the gouge after each entry and is immediately described and sampled, jointly by the archaeologist and the geomorphologist (Figure 4.4). Compared to the alternative method of cores extracted in sealed opaque plastic tubes and opened later in the laboratory, this one has the advantage of providing right away a view of the aspect and content of the cores (consistency and nature of sediments, presence of artefacts, charcoal, etc.), thus presenting the possibility to adjust the research strategy accordingly (e.g. continue deeper or stop, to repeat the drilling with different equipment, to take or not take samples for ¹⁴C dating). It is, however, less well adapted to a fine study of sensitive palaeobiological indicators (pollens, phytoliths, charcoal micro-remains, etc.). For this reason we also used sealed tubes for the lower part of one intra-site core (C9), as well as for two cores off-site (Dik5, Dik8) that were closer to the water source and therefore were susceptible to contain more such material, especially pollen.

The fact that the cores passed through the entire depth of the tell’s thick sedimentation allowed us to develop considerably our understanding of the types of sediments met, based on the prior good knowledge of the geomorphological context of the settlement, and on that of the anthropogenic layers acquired by the archaeologists during the recent excavation programs. We took also into account the guidelines provided in reference books for the detailed description of anthropogenic sediments on the basis of micromorphological analyses (Courty et al., 1989; Courty & Fedoroff, 2002). This combination of sedimentological and archaeological criteria allowed the distinction of several types of sediments (‘faciès’), seventeen in all (Table 4.1), ranging from those with little or no human presence (F1–F7) to those which are unquestionably anthropogenic (F8–F17). The number of types of sediment gives an idea of the variety of factors commanding the final aspect and consistency of the archaeological layers that form a tell.

Samples for ¹⁴C dating were taken from all layers that might be of interest for understanding the stratigraphical sequence, especially in the lowest parts. We preferred, of course, occupation layers in situ or only slightly disturbed, which were by chance those containing numerous organic remains, that is charcoal or, as an exception, seeds and bones. The samples were given for dating without any prior determination of species. This choice creates some problems for interpreting the results from charcoal samples, as we are unable to say whether the provided date concerns a short-lived organism and therefore contemporary with the observed sedimentary event, or a long-lived one that could have already been old when it was trapped in the sediment. We decided, however, to take the risk, in order to accelerate the dating process, as we estimated that the probability to fall constantly on parts from the innermost rings of multi-centenary oaks was altogether weak. In the off-site cores we also dated samples from the sediment itself (organic-rich), but these results are out of the scope of the present chapter and are not discussed here. The processing was made at the Centre for Radiocarbon Dating in Lyon (France) and the measurement at the Laboratory of Carbon-14 Measurement at the
Table 4.1: Types of Sediments Identified in the Dikili Tash Cores

<table>
<thead>
<tr>
<th>Type</th>
<th>Organisation</th>
<th>Sediments</th>
<th>Archaeological artefacts</th>
<th>Nature</th>
<th>Process of deposits</th>
<th>Post-depositional evolution</th>
<th>Interpretation</th>
</tr>
</thead>
<tbody>
<tr>
<td>F1</td>
<td>Coarse bedding</td>
<td>Sand and gravel</td>
<td>-</td>
<td>Channel deposits</td>
<td>Pedogenesis</td>
<td>Channel on the Pleistocene alluvial fan</td>
<td></td>
</tr>
<tr>
<td>F2</td>
<td>Coarse bedding</td>
<td>Yellowish sandy silts with gravel</td>
<td>-</td>
<td>Overbank deposits</td>
<td>Pedogenesis</td>
<td>Pleistocene alluvial fan</td>
<td></td>
</tr>
<tr>
<td>F3</td>
<td>Massif</td>
<td>Yellowish silt</td>
<td>-</td>
<td>Aeolian to colluvial</td>
<td>Pedogenesis</td>
<td>Aeolian silt reworked locally</td>
<td></td>
</tr>
<tr>
<td>F4</td>
<td>Aggregate</td>
<td>Brown silt</td>
<td>-</td>
<td>-</td>
<td>Pedogenesis</td>
<td>-</td>
<td></td>
</tr>
<tr>
<td>F5</td>
<td>Massif</td>
<td>Dark grey or brown silt</td>
<td>-</td>
<td>Runoff</td>
<td>Slight pedogenesis</td>
<td>Palustrine deposits</td>
<td></td>
</tr>
<tr>
<td>F6</td>
<td>Bedded</td>
<td>Oncolitic sand</td>
<td>+</td>
<td>Runoff</td>
<td>Slight pedogenesis</td>
<td>Carbonated alluvial deposits linked to the karstic stream</td>
<td></td>
</tr>
<tr>
<td>F7</td>
<td>Massif</td>
<td>Ochre silty sand</td>
<td>+</td>
<td>Particulate</td>
<td>-</td>
<td>Foxhole</td>
<td></td>
</tr>
<tr>
<td>F8</td>
<td>Thick horizon (&gt;20 cm) + aggregates</td>
<td>Dark brown silt</td>
<td>++</td>
<td>Charcoal, bone fragments, burnt clay</td>
<td>Erosion</td>
<td>Top of palaeosol with anthropogenic features</td>
<td></td>
</tr>
<tr>
<td>F9</td>
<td>Thin layer (&lt;10 cm) + aggregates</td>
<td>Dark brown silt</td>
<td>++</td>
<td>Charcoal, bone fragments, burnt clay</td>
<td>Erosion</td>
<td>Thin soil indicating an abandonment</td>
<td></td>
</tr>
<tr>
<td>F10</td>
<td>Random</td>
<td>Light brown sandy silt</td>
<td>++</td>
<td>Charcoal, bone fragments, burnt clay</td>
<td>Pedogenesis, erosion</td>
<td>Colluvial with pedogenic features</td>
<td></td>
</tr>
<tr>
<td>F11</td>
<td>Slightly organised</td>
<td>Light brown silt</td>
<td>+++</td>
<td>Charcoal, bone fragments, burnt clay, sherds</td>
<td>Pedogenesis, erosion</td>
<td>Reworked archaeological level in colluvial deposits</td>
<td></td>
</tr>
<tr>
<td>F12</td>
<td>Slightly organised</td>
<td>Yellowish brown-grey silt</td>
<td>++++</td>
<td>Charcoal, bone fragments, burnt clay, sherds</td>
<td>Pedogenesis, erosion</td>
<td>Slightly disturbed archaeological layer</td>
<td></td>
</tr>
<tr>
<td>F13</td>
<td>Slightly to strongly organised</td>
<td>Yellowish to reddish silt</td>
<td>++++</td>
<td>Construction earth slightly burnt</td>
<td>Archaeological structure reworked or in situ</td>
<td>Weathering</td>
<td>Mud structure or mud-brick, adobe</td>
</tr>
</tbody>
</table>
### Table 4.2 14C Dates From the 2012–2014 Cores at Dikili Tash (Calibration at 2σ, i.e. 95.4% Probability, With the IntCal13 Curve). Dates from the 2010 cores have been published in Lespez et al., 2013, and reproduced in this volume (Appendix), together with those from earlier cores A and B.

<table>
<thead>
<tr>
<th>Inv.no.</th>
<th>Sample</th>
<th>Exp. Lab.no.</th>
<th>Meas. Lab.no.</th>
<th>8P date ± 1σ BP</th>
<th>2σ BP</th>
<th>calBP date (2σ)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Core 7</td>
<td>1165–003 charcoal Lyon-10334 SacA-33943</td>
<td>6960</td>
<td>35</td>
<td>5971–5745</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Core 10</td>
<td>9245–005 charcoal Lyon-10333 SacA-33942</td>
<td>6875</td>
<td>30</td>
<td>5841–5707</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>9251–002 charcoal Lyon-10332 SacA-33941</td>
<td>7030</td>
<td>35</td>
<td>5996–5842</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>9255–002 charcoal Lyon-10331 SacA-33940</td>
<td>7150</td>
<td>30</td>
<td>6068–5986</td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>9262–001 charcoal Lyon-10330 SacA-33939</td>
<td>7200</td>
<td>30</td>
<td>6203–6003</td>
<td></td>
<td></td>
</tr>
<tr>
<td>Core 12</td>
<td>Dik12–221 charcoal Poz-65035 7530</td>
<td>50</td>
<td>6467–6254</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td></td>
<td>Dik12–233 charcoal Poz-65036 7585</td>
<td>50</td>
<td>6475–6405</td>
<td></td>
<td></td>
<td></td>
</tr>
</tbody>
</table>

Center of Atomic Energy at Saclay (France), with the AMS method; only two samples among those of interest to us here were dated at the Radiocarbon Laboratory of Poznań (Poland), again with AMS.

The succession of the different types of sediments in each core, combined with the radiocarbon dates obtained (Table 4.2), and their correlation, taking into account the compression effect (5–20 percent), enable us to interpret the formation process of the tell settlement.

**Interpretation of the Results and Consequences for our Knowledge about the Early Neolithic in the Area Under Study**

The results obtained so far allow us to reconstruct the evolution of settlement at Dikili Tash in a quite detailed manner, from the end of the Last Glacial period until the advanced stages of the Neolithic.

**The Pristine Soil**

The original soil was reached in all drillings: it consists of a light yellow, sandy or silty sediment (F1 to F3), whose formation dates back to the Pleistocene, on top of which developed during the early Holocene an argillaceous dark-coloured soil (F4); the latter is rich in organic materials, reflecting a slow decomposition in a forested environment (Figure 4.5). This was the soil first encountered by people that came to live at Dikili Tash, little after the mid-7th millennium BC (see below ii). Both the Pleistocene substratum and the Holocene palaeosol appear at slightly different altitudes, suggesting a gentle but clear north-west/south-east inclination, which corresponds to the natural slope of the large Pleistocene alluvial fan underlying the site's location. Such piedmont landscapes were probably forested since the beginning of the Holocene according to the pollen data of the marsh of Philippi (Wijmstra, 1969; Greig & Turner, 1974; Pross et al., 2009; Glais et al., 2015).

**Geometry of the Site and Position of the First Settlement (Figure 4.2, 6–8)**

Evidence of human activity – lithics alongside with tiny fragments of bones, amorphous burnt clay and charcoal – is present already in the topmost layers of the palaeosol in most of the intra-site cores (Lespez et al., 2013: 34–5 and Fig. 4). The date, however, of
Figure 4.5. Lowest Part of Core C10, Showing the Passage From Pleistocene Silt to the First Level of Occupation on Top of the Holocene Palaeosol

Figure 4.6. West–East Section of the Dikili Tash Tell With Presumed Evolution of the Settlement in the Different Periods. EN: Early Neolithic; MN: Middle Neolithic; LN1: Late Neolithic I; LN2: Late Neolithic II/Chalcolithic
these ‘intrusions’ is not the same everywhere: the earliest such remains are located indeed in the area of cores C2, C3 and C10, respectively in the eastern and northeastern parts of the future tell, which are also those closer to the water source. The available ^14^C dates here are for the years around 6400/6300 cal BC, that is practically one thousand years before the earliest layer excavated in sector 1, at the tell’s southern slope, and more than five hundred years before the earliest radiocarbon-dated deposits in the nearby core B of 1993 (supra). In the latter area, the beginning of human activity does not seem to go beyond 6100 or 6000 cal BC, as attested by a ^14^C date from the palaeosol’s top in core C1 (Lyon-7617). Therefore, it is not here that we should dig in order to unearth the earliest occupation remains at Dikili Tash. An even later date is obtained from the palaeosol’s top in core C4, to the northwest, indicating that this was the last part of the site to be visited by the Neolithic settlers. It should be recalled that a date around 6500/6400 cal BC (Lyon-5018) was also obtained from the Holocene palaeosol in the 1993 core A: initially left aside as an isolated poorly contextualised measurement, this date seems today perfectly coherent with the rest, indicating that the first human activities in the site developed along the eastern and northeastern foot of the present tell. Furthermore, one of the core drillings made in the off-site area at the bottom of the tell (Dik12, total depth 4 m) shows the development of a marshy palaeosol, which contains archaeological artefacts (but no architectural remains) and evidence of cereal cultivation (detailed presentation in Glais et al., 2015). Its dating from two radiocarbon dates to the middle of the 7th millennium BC (Poz-65035 and Poz-65036) indicates that the activities of the first settlers extended also in this area.

An occupation level in situ (F15) was found immediately on top of the palaeosol in the intra-site core C10 and another one, very slightly disturbed (F14), in C3. The former was rich in artefacts, especially sherds with monochrome (grey-black and reddish-brown) polished surfaces. Given the small distance between the two points, one could be tempted to assume that they belong to the same occupation episode, which would represent in this case the first settlement, properly speaking, at Dikili Tash. But the ^14^C date from C10 is younger than the one from C3 (~6200–6000 versus 6378–6220 cal BC, respectively Lyon-10330 and Lyon-7630). The explanation for this discrepancy seems to lie in the history of events in this area, in connection with water: the presence of a series of layers with silts and oncolithic sands (F5, F6) at less than 50 cm above the palaeosol in cores C2 and C3 suggests indeed a rise of the water level in the pond formed by the spring northeast of the site, dated to ~6200–6000 cal BC (Lyon-7625, Lyon-7629). The presence of marshy sediments near the base of C8 and C9 probably has to be connected with the same humid event, which would have affected these areas more severely. This incident probably stopped the development of the settlement at those precise spots, causing the inhabitants to move a few meters farther to the west, at C10. Building activity started again in the area of C2 and C3 only after the water level was back to normal and the area had filled with colluvia. No other events of that kind are seen in the following parts of the sequence.

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Figure 4.7. North-South Section of the Tell
Figure 4.8. Diagram With the 14C Dates From Cores at Dikili Tash (Intra-Site and Close Off-Site), Modelled With OxCal (v4.2.3) According to Their Attribution to the Successive Occupational Stages

The Settlement’s Evolution During the Next Centuries

The earliest layers in the cores taken from the other sides of the tell (C1, C7 and C5 to the south, C4 to the northwest) also date to the years after 6100/6000 and 5900/5800 cal BC, respectively, indicating a progressive expansion and at the same time a consolidation of the settlement at the onset of the 6th millennium BC. The succession of anthropogenic sediments in the different cores bear witness to the progressive construction of the tell. They show levels of occupation, abandonment and some thin layers of colluvial deposits. Nevertheless, the accumulation of anthropogenic deposits remained slow and Dikili Tash was a large flat site on the edge of a small marshy valley. The tell form was shaped in the last periods of the Neolithic, from 5400 BC onwards, through a succession of occupation levels in the same place as testified by archaeological excavations. The settlement was maintained more or less within the same limits during the rest of the 6th millennium.
and beyond, at least until the end of the 5th millennium bc, when it seems to have been abandoned for a few centuries (Darcque et al., 2014; Darcque et al., 2015: 410–4; Tsirtsoni, 2016).

CONCLUSIONS AND PERSPECTIVES

The discovery of Early and Middle Neolithic layers at the base of the tell at Dikili Tash puts an end to the long-lasting debate about the start of permanent settlements in eastern Macedonia and provides interesting clues concerning the processes of establishment of the newcomers under the new conditions (topographical, environmental, etc.). However, we cannot appreciate thoroughly the role and the dynamics of these conditions until those layers are properly excavated and studied, and not just looked at through a ‘keyhole’. Hence, this is in our plans for the following years.

According to the results from the cores, the best place to set the new excavation trench is the area around core C3, in the northeastern part of the tell: it was there, indeed, that the first practically undisrupted occupation levels were found directly on top of the Holocene palaeosol, rich in archaeological artefacts and radiocarbon-dated to 6400–6200 cal bc. The semi-circular end-scaper found at this spot is reminiscent of some Anatolian Pre-Pottery Neolithic B (PPNB) examples (see Lespez et al., 2013: 34). It also has parallels among lithic assemblages of the Marmara area, namely at Pendik, Barcm, Ilipnar and Aktopraklik (Gatsov, 2009, Fig. 49:3, 54:2; and personal information from I. Gatsov and P. Nedelcheva). Unfortunately, at this particular spot the thickness of later deposits is at least 8 m – meaning that we will need a heavy infrastructure and several years of excavation before getting to the base. A ‘shortcut’ could be offered by the close off-site core Dik12, where mid-7th millennium deposits are found at a depth of less than 3 m underneath later alluvial deposits. Excavation there would concern, however, only the earliest occupation settlement, and not the rest of the Early and Middle Neolithic sequence.

More generally, the results of the core-drilling program undertaken at Dikili Tash underline the benefits from a closer collaboration between archeologists and geomorphologists-sedimentologists. Such collaboration offers the possibility to reconstruct the complete geometry and history of an excavated archaeological site, to decide reliably about new areas to investigate, and to explore the archaeological potential of unexcavated sites or parts of sites. Naturally, a core-drilling program does not replace archaeological excavation (cf. Ammerman et al., 2008: 142). It remains difficult, indeed, to fully understand the sediment features due to the small quantities of material for observation, or to correlate sediments from distant drillings, as one has to constantly take into account the compression effects, the locally developed characteristics and of course the possibility of reworking. The task is even harder, when one has to deal with sedimentary units that are poor in datable material. Nevertheless, despite these difficulties, cores can securely testify earlier occupation layers hidden inside sites with thick stratification.

The methodology described presents a great potential for investigation in the area of eastern Macedonia and Thrace, with the particular characteristics of prehistoric settlements. Indeed, the experience from Dikili Tash shows that the absence of evidence for early settlements in such areas with great alluvial and/or anthropogenic sedimentation should not be taken as a fact, even when they are well investigated archaeologically. The absence of evidence is definitely not a reliable argument for the reconstruction of population movements and habitation practices in this area. In the next years we schedule a large-scale use of the method, and also a broadening of the issues treated through a multi-proxy analysis, in order to investigate the presence of anthropogenic remains at different spots, to understand the process of Neolithisation on a regional scale, and to establish the interconnection of human habitation with the local environment, and the development of this relation over time.

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